











Multi-approach gravity field models from Swarm GPS data

Signal and error in the Swarm models up to 2023-06-30

Delft University of Technology (TU Delft) Astronomical Institute of the University of Bern (AIUB) Astronomical Institute Ondřejov (ASU) Institute of Geodesy Graz (IfG) **Ohio State University (OSU)**

> Version 1.0 2023-10-01

Prepared and checked by João Encarnação Work Package Manager

Approved by Pieter Visser Project Manager 2023-10-01 Page 2 of 51

Contents

1	Vers	sion history	5
2	Intr	roduction	5
3	Sou	rce data	6
4	4.1	Chodology Combination	
5	Res 5.1 5.2 5.3 5.4 5.5 5.6	Spatial analysis	13 18 23 24
A	A.1 A.2 A.3	Astronomical Institute of the University of Bern	43 43
B Ra	B.1 B.2 B.3	Astronomical Institute of the University of Bern	45 45 45
N	Herei	ices	40
Li	ist o	of Figures	
	1 2 3 4	Monthly Satellite Laser Ranging (SLR)-derived C_{20} from Loomis and Rachlin (2020) (TN-14), compared to Cheng and Ries (2018) (TN-11) and Cheng and Ries (2019) (CSR-RL06, from Gravity Recovery And Climate Experiment (GRACE)/GRACE Follow On (GRACE-FO))	77 88 9

2023-10-01 Page 3 of 51

5	Epoch-wise cumulative spatial RMS (top) and its global average (bottom) of
	the difference between Swarm GFMs and GRACE, over land areas, considering
	750km Gaussian smoothing. This is (an estimate of) the evolution of the ability
	of the various Swarm solutions to predict land mass transport processes over
	time (top) and its global sum (bottom)
6	Epoch-wise cumulative spatial RMS (top) and its global sum (bottom) of the
	difference between Swarm GFMs and GRACE, over ocean areas, considering
	750km Gaussian smoothing. This is the epoch-wise quality of the Swarm GFMs,
	and reported in the header of the combined GFMs files. The Swarm combined
	model (RL01) includes the new version of the IfG and ASU individual solutions
	only after December 2021; this is the reason why the combined model is not
	below these individual solutions before 2020
7	Per-degree mean (top) and its overall cumulative (bottom) of the correlation
	coefficient between Swarm GFMs and GRACE, over land areas, considering
	750km Gaussian smoothing. The temporal correlation at every Stokes coeffi-
	cient is computed and the average over each degree is plotted at the top. It
	illustrates how well the temporal variations of the Swarm models agree with
	what is predicted from the GRACE/GRACE-FO climatological model
8	Per-degree mean (top) and its overall cumulative (bottom) of the correlation
U	coefficient between Swarm GFMs and GRACE, over ocean areas, considering
	750km Gaussian smoothing. It illustrates that the Swarm models agree poorly
	with the mass variations over the ocean as predicted by the GRACE/GRACE-FO
	climatological model
9	Per-degree mean (top) and its overall cumulative (bottom) of the correlation
J	coefficient between Swarm and GRACE/GRACE-FO GFMs (not the GRACE/
	GRACE-FO climatological model), globally and with no smoothing. It illus-
	trates that the Swarm models fail to represent the same temporal variations as
	GRACE/GRACE-FO above degree 15-20
10	Per-coefficient RMS difference between Swarm GFMs and GRACE considering
10	· ·
	750km Gaussian smoothing, over land (left column) and ocean (right column)
	areas, for AIUB, ASU, IfG, OSU and combined solutions (respectively from top
11	to bottom)
11	Per-coefficient correlation coefficient between Swarm GFMs and GRACE con-
	sidering 750km Gaussian smoothing, over land (left column) and ocean (right
	column) areas, for AIUB, ASU, IfG, OSU and combined solutions (respectively
10	from top to bottom)
12	Monthly degree-RMS for the 3 most recent months, all individual and combined
10	Swarm solutions, as well as GRACE/GRACE-FO (no smoothing)
13	Time series of EWH for the Amazon basin (latitude -17 to 3 degrees, longitude
	-76 to -47 degrees)
14	Time series of EWH for the Orinoco basin (latitude -3 to 12 degrees, longitude
	-72 to -59 degrees)
15	Time series of EWH for the La Plata basin (latitude -34 to -19 degrees, longitude
	-65 to -50 degrees)
16	Time series of EWH for the Mississippi basin (latitude 29 to 44 degrees, longitude
	-101 to -80 degrees)
17	Time series of EWH for the Columbia region (latitude 38 to 50 degrees, longitude
	-125 to -110 degrees)

2023-10-01 Page 4 of 51

18	Time series of EWH for the Alaska (latitude 56 to 65 degrees, longitude -151 to -129 degrees)	29
19	Time series of EWH for the Western Greenland region (latitude 60 to 85 degrees,	20
	longitude -60 to -37 degrees)	30
20	Time series of EWH for the Danube basin (latitude 43 to 48 degrees, longitude 13 to 28 degrees)	31
21	Time series of EWH for the Western Sub-Saharan basin (latitude 5 to 15 degrees,	31
21	longitude -15 to -1 degrees)	32
22	Time series of EWH for the Eastern Sub-Saharan basin (latitude 1 to 13 degrees,	32
22	longitude -8 to 35 degrees)	33
23	Time series of EWH for the Congo and Zambezi basins (latitude -23 to -3 degrees,	55
23	longitude 14 to 38 degrees)	34
24	Time series of EWH for the Volga basin (latitude 53 to 61 degrees, longitude 34	54
2-1	to 56 degrees)	35
25	Time series of EWH for the Siberia region (latitude 57 to 72 degrees, longitude	55
23	68 to 109 degrees)	36
26	Time series of EWH for the Ganges-Brahmaputra basin (latitude 15 to 30 degrees,	00
_0	longitude 72 to 89 degrees)	37
27	Time series of EWH for the Indochina region (latitude 12 to 29 degrees, longitude	٥.
	93 to 105 degrees)	38
28	Time series of EWH for the Northern Australia region (latitude -24 to -10 degrees,	00
	longitude 124 to 145 degrees)	39
29	Time series of EWH for the Western Antarctica region (latitude -80 to -70 degrees,	
	longitude -140 to -85 degrees)	40
30	Time series of EWH for the Eastern Antarctica region (latitude -80 to -68 degrees,	
	longitude 80 to 130 degrees)	41
31	Temporal variability of the Swarm combined solutions	42
T ! - 4 -	£75-1.1	
List o	of Tables	
1	Overview of the gravity field estimation approaches	6
2	Versions of the GFMs, and the Kinematic Orbits (KOs) used in their estimation,	
	relevant to this report	6
3	Statistics of the agreement between GRACE/GRACE-FO and Swarm time series	
	relative to the GRACE/GRACE-FO climatological model for the Amazon basin. $$.	24
4	Statistics of the agreement between GRACE/GRACE-FO and Swarm time series	
	$relative \ to \ the \ GRACE/GRACE-FO \ climatological \ model \ for \ the \ Orino co \ basin. \ .$	25
5	Statistics of the agreement between GRACE/GRACE-FO and Swarm time series	
	relative to the GRACE/GRACE-FO climatological model for the La Plata basin. $$.	26
6	Statistics of the agreement between GRACE/GRACE-FO and Swarm time series	
	relative to the GRACE/GRACE-FO climatological model for the Mississippi basin.	27
7	Statistics of the agreement between GRACE/GRACE-FO and Swarm time series	
	relative to the GRACE/GRACE-FO climatological model for the Columbia region.	28
8	Statistics of the agreement between GRACE/GRACE-FO and Swarm time series	
_	relative to the GRACE/GRACE-FO climatological model for the Alaska	29
9	Statistics of the agreement between GRACE/GRACE-FO and Swarm time series	
	relative to the GRACE/GRACE-FO climatological model for the Western Green-	0.0
	land region	30

2023-10-01 Page 5 of 51

10	Statistics of the agreement between GRACE/GRACE-FO and Swarm time series relative to the GRACE/GRACE-FO climatological model for the Danube basin 3
11	Statistics of the agreement between GRACE/GRACE-FO and Swarm time series
	relative to the GRACE/GRACE-FO climatological model for the Western Sub-
	Saharan basin
12	Statistics of the agreement between GRACE/GRACE-FO and Swarm time series relative to the GRACE/GRACE-FO climatological model for the Eastern Sub-
	Saharan basin
13	Statistics of the agreement between GRACE/GRACE-FO and Swarm time series relative to the GRACE/GRACE-FO climatological model for the Congo and Zam-
	bezi basins
14	Statistics of the agreement between GRACE/GRACE-FO and Swarm time series
	relative to the GRACE/GRACE-FO climatological model for the Volga basin 3
15	Statistics of the agreement between GRACE/GRACE-FO and Swarm time series
	relative to the GRACE/GRACE-FO climatological model for the Siberia region 3
16	Statistics of the agreement between GRACE/GRACE-FO and Swarm time series
	$relative \ to \ the \ GRACE/GRACE-FO\ climatological\ model for \ the\ Ganges-Brahmaputra$
	basin
17	Statistics of the agreement between GRACE/GRACE-FO and Swarm time series
	relative to the GRACE/GRACE-FO climatological model for the Indochina region. 3
18	Statistics of the agreement between GRACE/GRACE-FO and Swarm time series
	relative to the GRACE/GRACE-FO climatological model for the Northern Aus-
	tralia region
19	Statistics of the agreement between GRACE/GRACE-FO and Swarm time series
	relative to the GRACE/GRACE-FO climatological model for the Western Antarc-
20	tica region
20	Statistics of the agreement between GRACE/GRACE-FO and Swarm time series
	relative to the GRACE/GRACE-FO climatological model for the Eastern Antarctics region
21	tica region
21	regions displayed in Sections 5.5.1 to 5.5.18
	regions displayed in sections 3.3.1 to 3.3.18

1 Version history

Version 1, 2023-10-01

• Validation of combined models version 09, from start of mission until 2023-06-30.

2 Introduction

We report some statistics of the individual and combined GFMs produced on the context of the *Multi-approach gravity field models from Swarm GPS data* project. The approach for combining individual gravity field solutions, i.e. those produced by the various partners mentioned in Section 3, is described in Section 4.1. The procedure and assumption used to derive the statistics is described in Section 4.2. Finally, the results are presented in Section 5.

This report does not intend to draw conclusions regarding the presented statistics, it is merely a descriptive document of the signal and error in the individual and combined Swarm

2023-10-01 Page 6 of 51

GFMs. For this reason, the text in Section 5 is restricted to clarifying the quantities shown in the plots.

3 Source data

The individual gravity field solutions are produced by the institutes listed in Table 1.

Table 1 – Overview of the gravity field estimation approaches

Inst.	Approach	Reference
AIUB	Celestial Mechanics Approach	Jäggi et al. (2016)
ASU	Decorrelated Acceleration Approach	Bezděk et al. (2016)
IfG	Short-Arcs Approach	Suesser-Rechberger et al. (2022)
OSU	Improved Energy Balance Approach	Guo et al. (2015)

Additional details about the different gravity field approaches can be found in (Teixeira da Encarnação and Visser, 2017).

The version of the individual GFMs is listed in Table 2.

Table 2 – Versions of the GFMs, and the KOs used in their estimation, relevant to this report.

Gravity Field Model	version	Kinematic Orbit
AIUB	01	AIUB
ASU	02 - 03	IfG
IfG	03 - 07	IfG
OSU	02	AIUB
combined	09	N/A

The version numbers listed in Table 2 are relevant within the project and are reported so that it is possible to trace back the results presented in Section 5. Particular to the combined models, version 09 relates to the chosen combination strategy, as concluded from Teixeira da Encarnação and Visser (2019).

4 Methodology

4.1 Combination

The combination of the models is conducted at the level of the solutions considering weights derived from Variance Component Estimation (VCE). As demonstrated in Teixeira da Encarnação and Visser (2019), the combination at the level of Normal Equation (NEQ) disagreed more with GRACE/GRACE-FO, as a result of the vastly different amplitudes of formal errors.

The combination considers the complete degree range (degrees 2 to 40) but the VCE weights are derived from degrees 2-20. This approach addresses the very high errors above degree 20, which would otherwise drive the value of the weights.

It is feasible to determine the VCE weights because there are two time-series based on AIUB orbits (i.e. AIUB and OSU) and two time-series based on IfG orbits (i.e. IfG and ASU). Therefore the impact of the KOs on the solutions and on the VCE weights is balanced.

2023-10-01 Page 7 of 51

4.2 Validation

The validation is done by comparing the individual and combined solutions to the Release 6 (RL06) GRACE/GRACE-FO GFMs produced at Center for Space Research (CSR), considering all solutions available at the this document is produced.

All solutions undergo a 750km radius spherical cap Gaussian filtering, unless otherwise noted, to clearly show the geophysical signal contained in the Swarm solutions. The GRACE and GOCE Gravity Model 05 (GGM05G) (Bettadpur et al., 2015) static GFM is subtracted from all models in order to isolate the time-variable component of Earth's gravity field. We chose to show the gravity field in terms of EWH, except for the statistics related to the correlation coefficient, which are non-dimensional as usual. The GRACE/GRACE-FO gravity field time series is linearly interpolated to the mid-month epoch of the Swarm solutions. The GRACE/GRACE-FO climatological model is evaluated at the same time domain. The analysis spans 2016-01-01 until 2023-06-30.

Note that there is no effort to meticulously consider or implement proper leakage reduction methods, e.g. by Guo, Duan and Shum (2010), in any of our analyses.

4.2.1 Earth's oblateness

The $C_{2,0}$ coefficient in all solutions has been replaced by the time series provided in Loomis and Rachlin (2020).

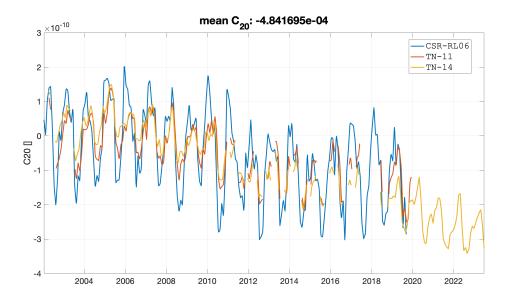


Figure 1 – Monthly SLR-derived C_{20} from Loomis and Rachlin (2020) (TN-14), compared to Cheng and Ries (2018) (TN-11) and Cheng and Ries (2019) (CSR-RL06, from GRACE/GRACE-FO).

4.2.2 Land and ocean analyses

Some analyses are restricted to either the land or ocean areas. In those cases, the land or ocean mask is applied in the spatial domain and a Spherical Harmonic (SH) analysis is done on the masked grid. The ocean mask excludes the coastal ocean areas that are roughly 1000km or less from land areas, as shown in Figure 2, while the land mask has no buffer zone.

2023-10-01 Page 8 of 51

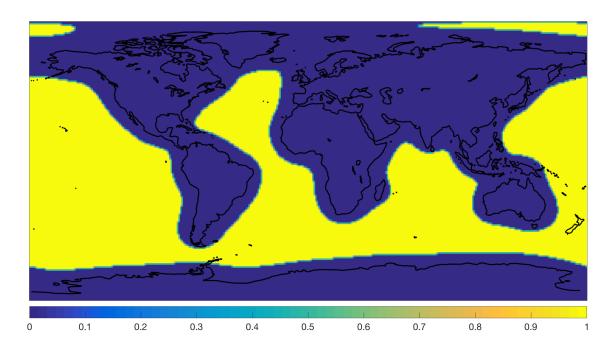


Figure 2 - Deep ocean mask.

4.2.3 Catchment time series

In Section 5.5, the geophysical signal represented by the Swarm solutions is evaluated on the basis of the time series of average EWH over restricted geographical locations, shown in Figure 3. Each averaging is done over the corresponding spatial truncation of an equiangular grid representation of the SH coefficients. The locations shown in Sections 5.5.1 to 5.5.18 are related to the largest hydrological basins and polar regions with the highest signal variability observed by GRACE/GRACE-FO. We perform a parametric regression on the time series of all geographical regions considering a constant and drift terms, along with annual and semiannual sine and co-sine terms to improve the robustness. We plot the linear part of this regression, in order to quantify the accuracy of Swarm-derived climatological trends. The time series are plotted along with tables presenting some statistics. The values of the constant and linear terms for the Swarm and GRACE/GRACE-FO solutions (column 1) are show in terms of EWH (columns 2 and 4). Additionally, the difference of these parameters between the Swarm and GRACE/GRACE-FO solutions relative to the GRACE/GRACE-FO climatological model is listed in columns 3 and 5 (the values for the latter data set in these columns is zero). Finally, the correlation coefficients is presented in the last column (the value for GRACE/GRACE-FO climatological model is 1). The constant term is the average basin storage over the relevant data period.

Page 9 of 51

temporal STD of GRACE (2014-01 to 2023-06) 750km Gaussian smoothing

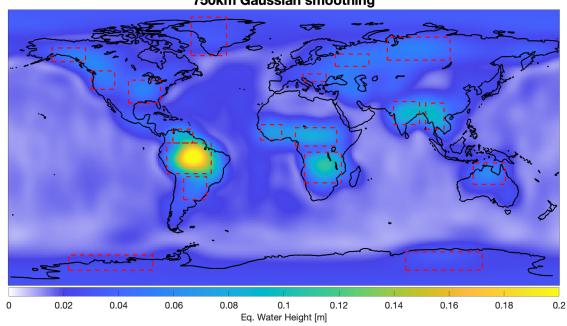
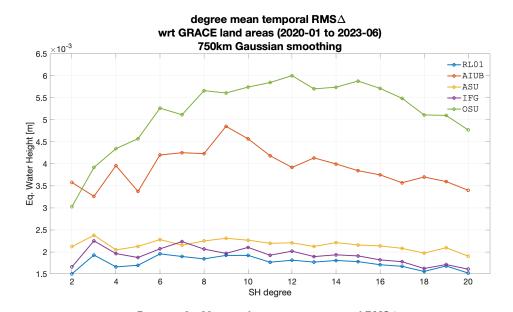


Figure 3 - Temporal variability of GRACE/GRACE-FO, including the boundaries of the regions analysed in Sections 5.5.1 to 5.5.18.

5 Results

5.1 Spatial analysis

5.1.1 Degree-mean RMS difference over land



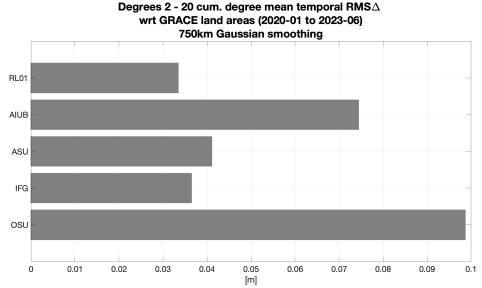
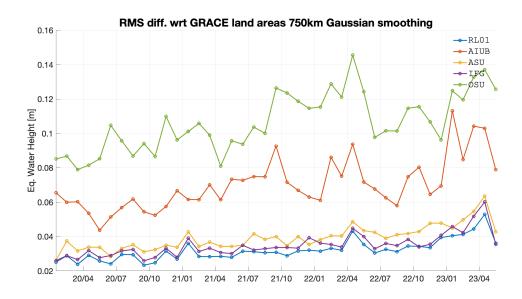


Figure 4 – Per-degree mean of the RMS difference (top) and cumulative degree-mean temporal RMS difference (bottom) between the Swarm GFMs and GRACE, over land areas, considering 750km Gaussian smoothing. This is (an estimate of) the average per-degree quality of the various Swarm solutions in the spectral domain (top) and globally (bottom). The degree amplitudes remain relatively constant with increasing degree, instead of growing in terms of EWH, as the result of the smoothing.

5.1.2 Cumulative degree amplitude difference over land



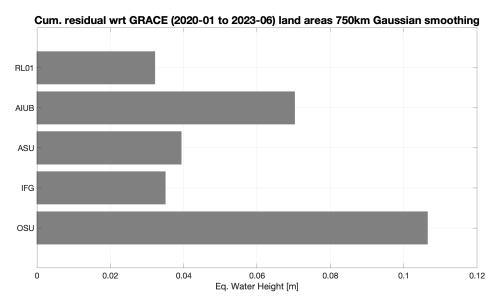
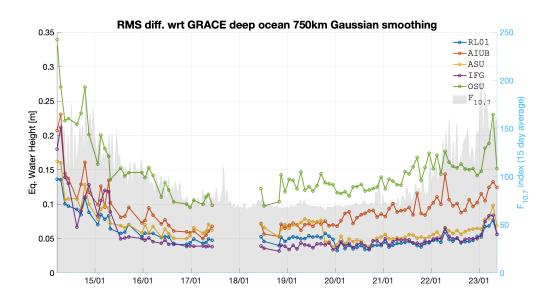


Figure 5 – Epoch-wise cumulative spatial RMS (top) and its global average (bottom) of the difference between Swarm GFMs and GRACE, over land areas, considering 750km Gaussian smoothing. This is (an estimate of) the evolution of the ability of the various Swarm solutions to predict land mass transport processes over time (top) and its global sum (bottom).

5.1.3 Cumulative degree amplitude difference over oceans



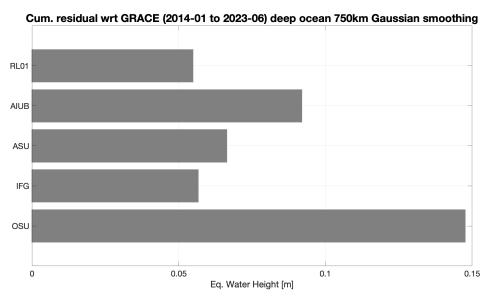
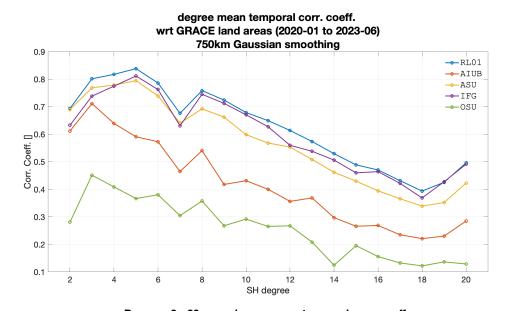


Figure 6 – Epoch-wise cumulative spatial RMS (top) and its global sum (bottom) of the difference between Swarm GFMs and GRACE, over ocean areas, considering 750km Gaussian smoothing. This is the epoch-wise quality of the Swarm GFMs, and reported in the header of the combined GFMs files. The Swarm combined model (RL01) includes the new version of the IfG and ASU individual solutions only after December 2021; this is the reason why the combined model is not below these individual solutions before 2020.

2023-10-01 Page 13 of 51

5.2 Temporal analysis

5.2.1 Per-degree mean correlation coefficient over land



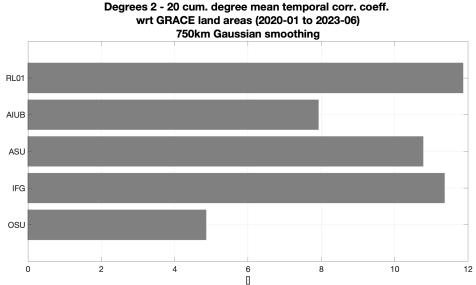
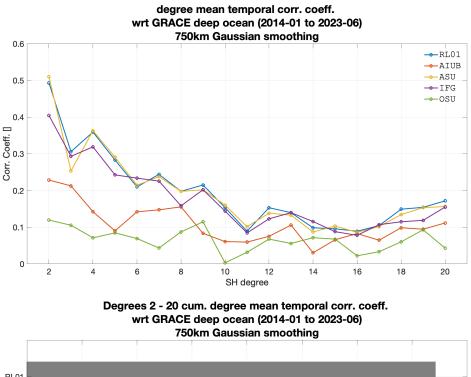


Figure 7 – Per-degree mean (top) and its overall cumulative (bottom) of the correlation coefficient between Swarm GFMs and GRACE, over land areas, considering 750km Gaussian smoothing. The temporal correlation at every Stokes coefficient is computed and the average over each degree is plotted at the top. It illustrates how well the temporal variations of the Swarm models agree with what is predicted from the GRACE/GRACE-FO climatological model.

5.2.2 Per-degree mean correlation coefficient over oceans



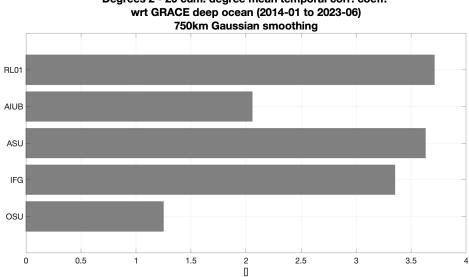


Figure 8 – Per-degree mean (top) and its overall cumulative (bottom) of the correlation coefficient between Swarm GFMs and GRACE, over ocean areas, considering 750km Gaussian smoothing. It illustrates that the Swarm models agree poorly with the mass variations over the ocean as predicted by the GRACE/GRACE-FO climatological model.

5.2.3 Global unsmoothed per-degree mean correlation coefficient

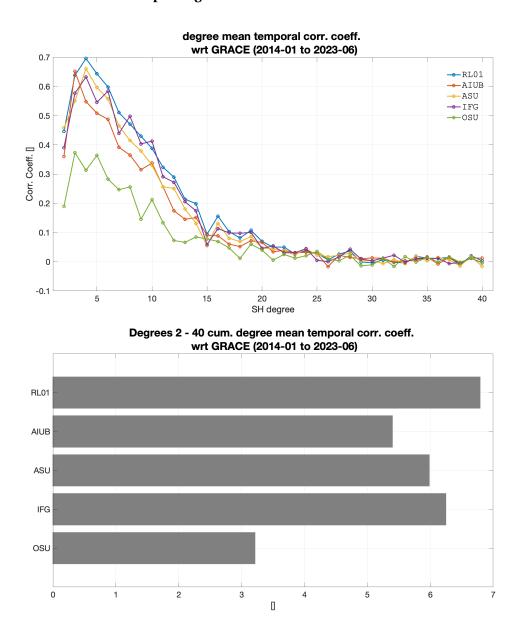


Figure 9 – Per-degree mean (top) and its overall cumulative (bottom) of the correlation coefficient between Swarm and GRACE/GRACE-FO GFMs (not the GRACE/GRACE-FO climatological model), globally and with no smoothing. It illustrates that the Swarm models fail to represent the same temporal variations as GRACE/GRACE-FO above degree 15-20.

2023-10-01 Page 16 of 51

5.2.4 Triangular plots of the RMS differences

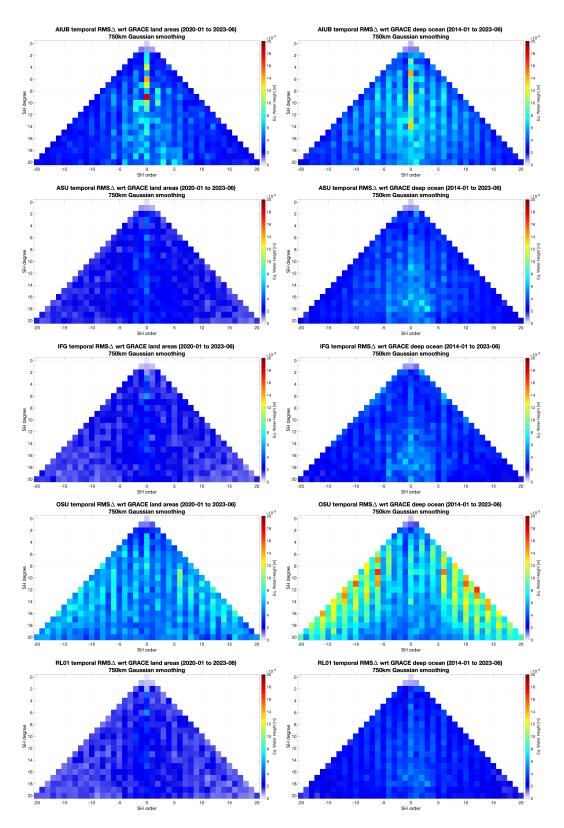


Figure 10 – Per-coefficient RMS difference between Swarm GFMs and GRACE considering 750km Gaussian smoothing, over land (left column) and ocean (right column) areas, for AIUB, ASU, IfG, OSU and combined solutions (respectively from top to bottom).

2023-10-01 Page 17 of 51

5.2.5 Triangular plots of the correlation coefficients

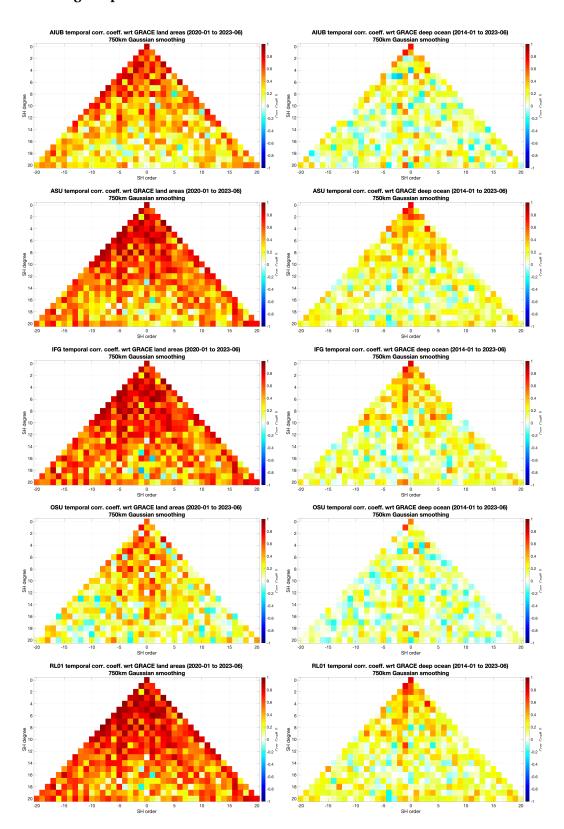
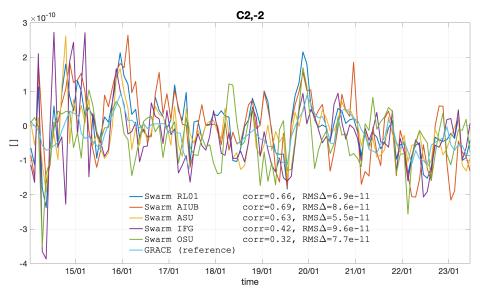


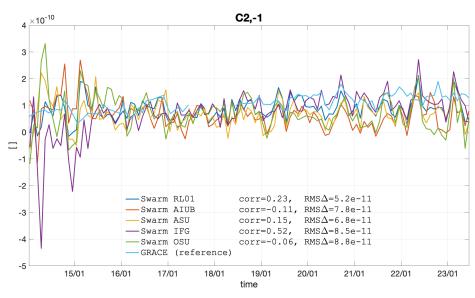
Figure 11 – Per-coefficient correlation coefficient between Swarm GFMs and GRACE considering 750km Gaussian smoothing, over land (left column) and ocean (right column) areas, for AIUB, ASU, IfG, OSU and combined solutions (respectively from top to bottom).

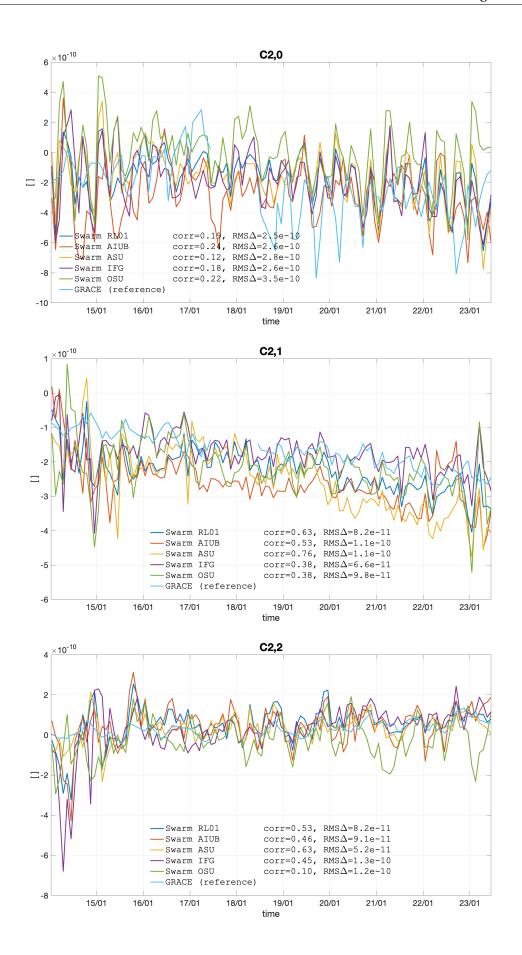
2023-10-01 Page 18 of 51

5.3 Low-degree zonal coefficients

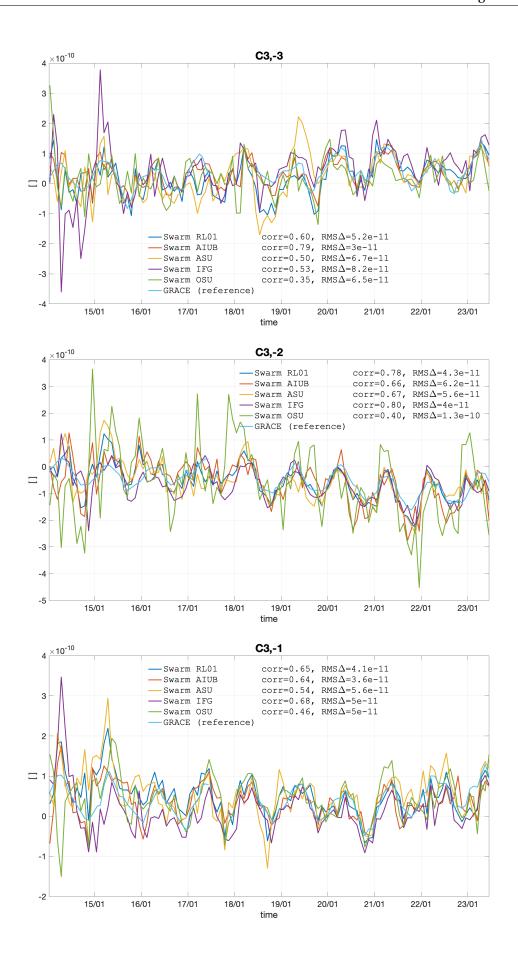
This sections presents the time series of the degree 2 and 3 coefficients (for all orders), showing their values in the Swarm and GRACE/GRACE-FO GFMs. The statistics in the legend consider GRACE as reference.



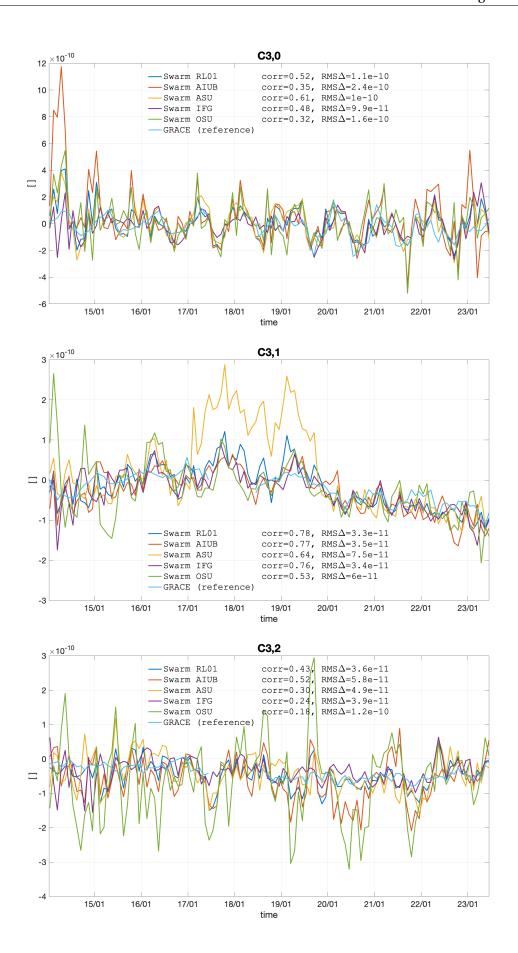




2023-10-01 Page 20 of 51

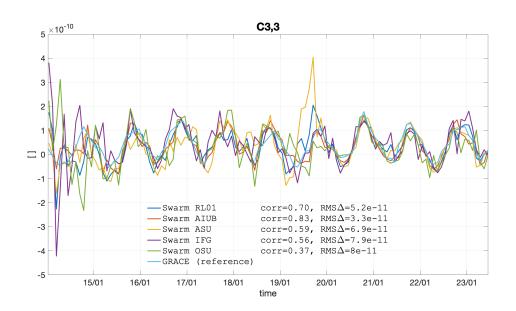


2023-10-01 Page 21 of 51



Multi-approach gravity field models from Swarm GPS data SW_VR_DUT_GS_0018 version 1.0

2023-10-01 Page 22 of 51



5.4 Monthly models

5.4.1 Monthly degree-RMS

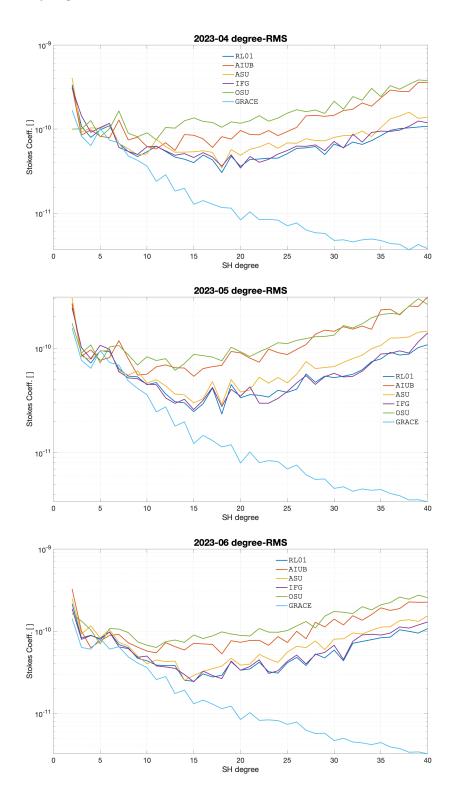


Figure 12 – Monthly degree-RMS for the 3 most recent months, all individual and combined Swarm solutions, as well as GRACE/GRACE-FO (no smoothing).

5.5 Time series of storage catchments

5.5.1 Amazon basin

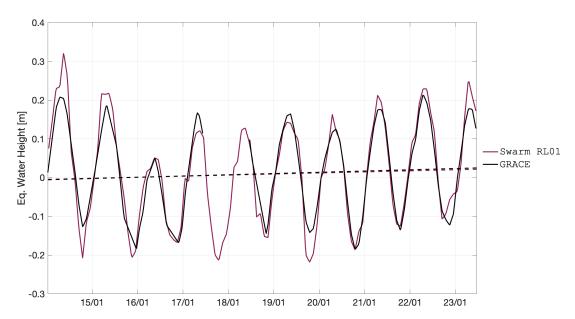


Figure 13 – Time series of EWH for the Amazon basin (latitude -17 to 3 degrees, longitude -76 to -47 degrees).

solution	constant term [cm]	constant term Δ [cm]	linear term [cm/year]	linear term Δ [cm/year]	corr. coeff.
Swarm RL01	1.41	-1.04	0.27	-0.05	0.96
GRACE	2.45	0.00	0.32	0.00	1.00

Table 3 – Statistics of the agreement between GRACE/GRACE-FO and Swarm time series relative to the GRACE/GRACE-FO climatological model for the Amazon basin.

5.5.2 Orinoco basin

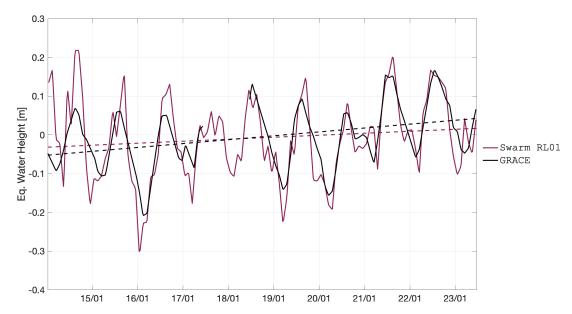


Figure 14 – Time series of EWH for the Orinoco basin (latitude -3 to 12 degrees, longitude -72 to -59 degrees).

solution	constant	constant	linear term	linear term	corr. coeff.
solution	term [cm]	term Δ [cm]	[cm/year]	Δ [cm/year]	[]
Swarm RL01	-1.03	-0.41	0.51	-0.49	0.85
GRACE	-0.62	0.00	1.00	0.00	1.00

Table 4 – Statistics of the agreement between GRACE/GRACE-FO and Swarm time series relative to the GRACE/GRACE-FO climatological model for the Orinoco basin.

5.5.3 La Plata basin

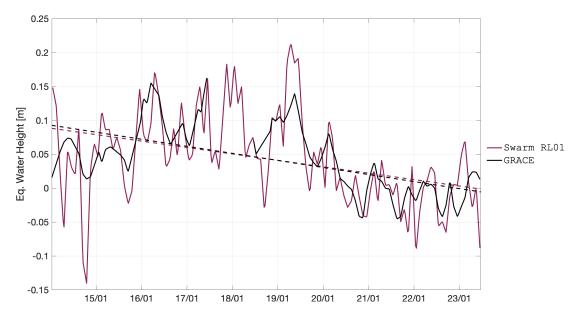


Figure 15 – Time series of EWH for the La Plata basin (latitude -34 to -19 degrees, longitude -65 to -50 degrees).

solution	constant	constant	linear term	linear term	corr. coeff.
solution	term [cm]	term Δ [cm]	[cm/year]	Δ [cm/year]	[]
Swarm RL01	4.46	0.28	-0.95	0.09	0.76
GRACE	4.18	0.00	-1.04	0.00	1.00

Table 5 – Statistics of the agreement between GRACE/GRACE-FO and Swarm time series relative to the GRACE/GRACE-FO climatological model for the La Plata basin.

5.5.4 Mississippi basin

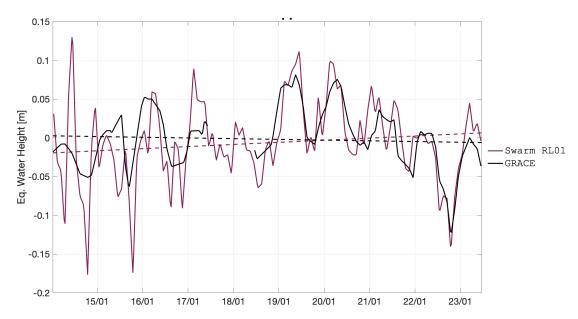


Figure 16 – Time series of EWH for the Mississippi basin (latitude 29 to 44 degrees, longitude -101 to -80 degrees).

solution	constant	constant	linear term	linear term	corr. coeff.
Solution	term [cm]	term Δ [cm]	[cm/year]	Δ [cm/year]	[]
Swarm RL01	-0.50	-0.62	0.27	0.36	0.70
GRACE	0.13	0.00	-0.09	0.00	1.00

Table 6 – Statistics of the agreement between GRACE/GRACE-FO and Swarm time series relative to the GRACE/GRACE-FO climatological model for the Mississippi basin.

5.5.5 Columbia region

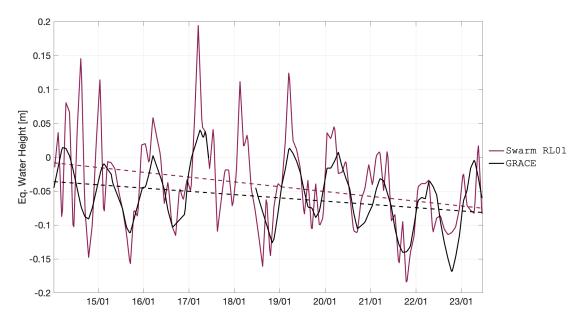


Figure 17 – Time series of EWH for the Columbia region (latitude 38 to 50 degrees, longitude -125 to -110 degrees).

solution	constant	constant	linear term	linear term	corr. coeff.
Solution	term [cm]	term Δ [cm]	[cm/year]	Δ [cm/year]	[]
Swarm RL01	-3.95	1.59	-0.71	-0.23	0.68
GRACE	-5.54	0.00	-0.48	0.00	1.00

Table 7 – Statistics of the agreement between GRACE/GRACE-FO and Swarm time series relative to the GRACE/GRACE-FO climatological model for the Columbia region.

2023-10-01 Page 29 of 51

5.5.6 Alaska

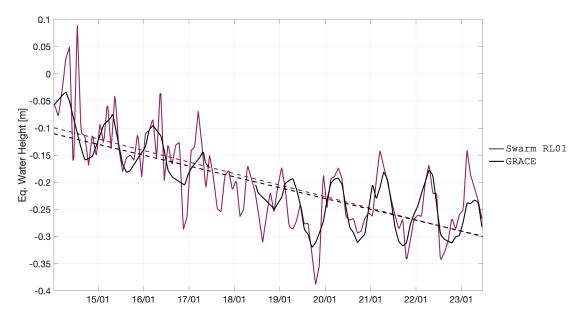


Figure 18 – Time series of EWH for the Alaska (latitude 56 to 65 degrees, longitude -151 to -129 degrees).

solution	constant term [cm]	constant term Δ [cm]	linear term [cm/year]	linear term Δ [cm/year]	corr. coeff.
Swarm RL01	-19.80	1.02	-2.13	-0.13	0.87
GRACE	-20.81	0.00	-2.00	0.00	1.00

Table 8 – Statistics of the agreement between GRACE/GRACE-FO and Swarm time series relative to the GRACE/GRACE-FO climatological model for the Alaska.

5.5.7 Western Greenland region

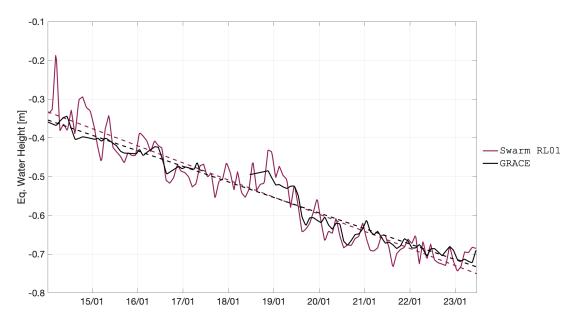


Figure 19 – Time series of EWH for the Western Greenland region (latitude 60 to 85 degrees, longitude -60 to -37 degrees).

solution	constant	constant	linear term	linear term	corr. coeff.
	term [cm]	term Δ [cm]	[cm/year]	Δ [cm/year]	[]
Swarm RL01	-54.14	1.80	-4.42	-0.41	0.96
GRACE	-55.95	0.00	-4.01	0.00	1.00

Table 9 – Statistics of the agreement between GRACE/GRACE-FO and Swarm time series relative to the GRACE/GRACE-FO climatological model for the Western Greenland region.

2023-10-01 Page 31 of 51

5.5.8 Danube basin

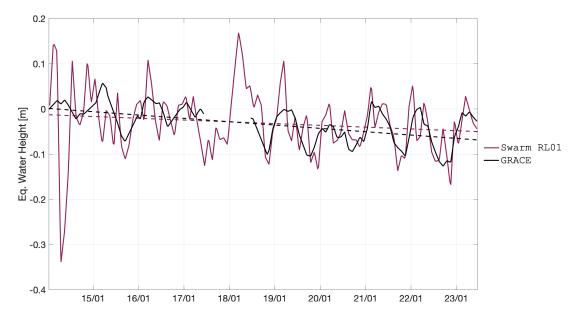


Figure 20 – Time series of EWH for the Danube basin (latitude 43 to 48 degrees, longitude 13 to 28 degrees).

solution	constant	constant	linear term	linear term	corr. coeff.
	term [cm]	term Δ [cm]	[cm/year]	Δ [cm/year]	[]
Swarm RL01	-3.00	0.27	-0.39	0.34	0.36
GRACE	-3.26	0.00	-0.74	0.00	1.00

Table 10 - Statistics of the agreement between GRACE/GRACE-FO and Swarm time series relative to the GRACE/GRACE-FO climatological model for the Danube basin.

5.5.9 Western Sub-Saharan basin

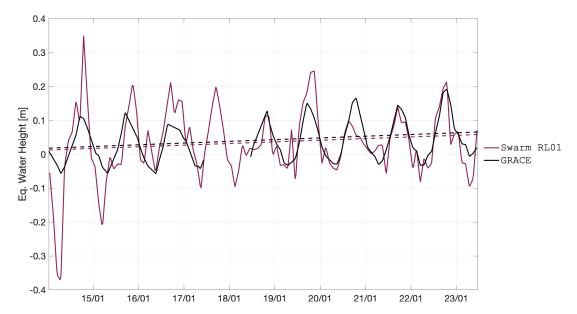


Figure 21 – Time series of EWH for the Western Sub-Saharan basin (latitude 5 to 15 degrees, longitude -15 to -1 degrees).

solution	constant	constant	linear term	linear term	corr. coeff.
	term [cm]	term Δ [cm]	[cm/year]	Δ [cm/year]	[]
Swarm RL01	3.11	-0.43	0.48	-0.02	0.72
GRACE	3.54	0.00	0.51	0.00	1.00

Table 11 - Statistics of the agreement between GRACE/GRACE-FO and Swarm time series relative to the GRACE/GRACE-FO climatological model for the Western Sub-Saharan basin.

5.5.10 Eastern Sub-Saharan basin

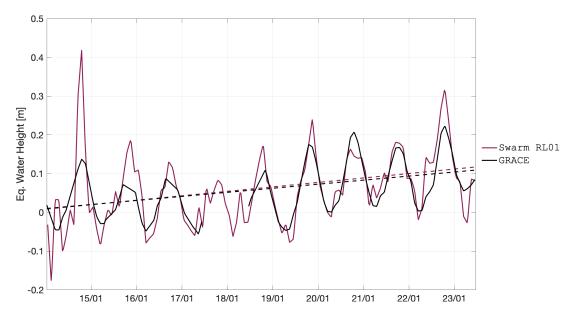


Figure 22 – Time series of EWH for the Eastern Sub-Saharan basin (latitude 1 to 13 degrees, longitude -8 to 35 degrees).

solution	constant	constant	linear term	linear term	corr. coeff.
	term [cm]	term Δ [cm]	[cm/year]	Δ [cm/year]	[]
Swarm RL01	5.93	0.34	1.15	0.10	0.83
GRACE	5.60	0.00	1.05	0.00	1.00

Table 12 – Statistics of the agreement between GRACE/GRACE-FO and Swarm time series relative to the GRACE/GRACE-FO climatological model for the Eastern Sub-Saharan basin.

5.5.11 Congo and Zambezi basins

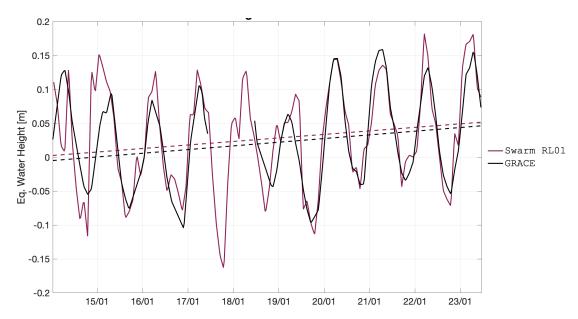


Figure 23 – Time series of EWH for the Congo and Zambezi basins (latitude -23 to -3 degrees, longitude 14 to 38 degrees).

solution	constant	constant	linear term	linear term	corr. coeff.
	term [cm]	term Δ [cm]	[cm/year]	Δ [cm/year]	[]
Swarm RL01	3.06	-0.07	0.52	-0.02	0.84
GRACE	3.13	0.00	0.54	0.00	1.00

Table 13 – Statistics of the agreement between GRACE/GRACE-FO and Swarm time series relative to the GRACE/GRACE-FO climatological model for the Congo and Zambezi basins.

5.5.12 Volga basin

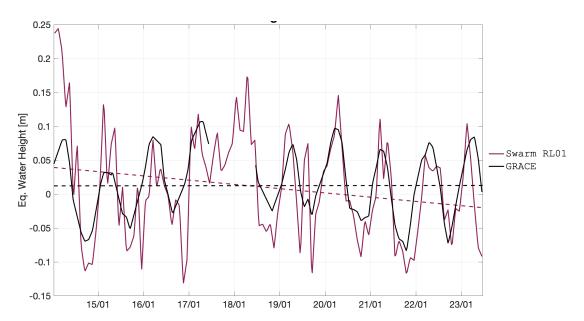


Figure 24 – Time series of EWH for the Volga basin (latitude 53 to 61 degrees, longitude 34 to 56 degrees).

solution	constant	constant	linear term	linear term	corr. coeff.
	term [cm]	term Δ [cm]	[cm/year]	Δ [cm/year]	[]
Swarm RL01	1.25	-0.61	-0.63	-0.63	0.73
GRACE	1.86	0.00	0.01	0.00	1.00

Table 14 – Statistics of the agreement between GRACE/GRACE-FO and Swarm time series relative to the GRACE/GRACE-FO climatological model for the Volga basin.

5.5.13 Siberia region

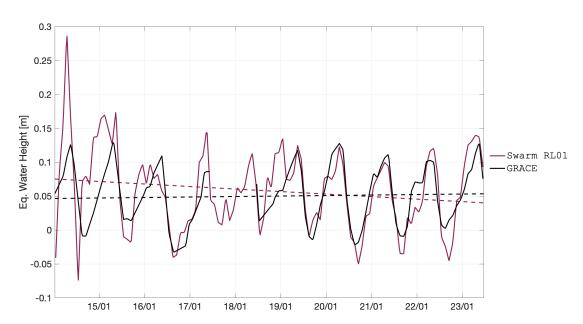


Figure 25 – Time series of EWH for the Siberia region (latitude 57 to 72 degrees, longitude 68 to 109 degrees).

solution	constant	constant	linear term	linear term	corr. coeff.
	term [cm]	term Δ [cm]	[cm/year]	Δ [cm/year]	[]
Swarm RL01	5.98	0.41	-0.37	-0.45	0.75
GRACE	5.57	0.00	0.07	0.00	1.00

Table 15 – Statistics of the agreement between GRACE/GRACE-FO and Swarm time series relative to the GRACE/GRACE-FO climatological model for the Siberia region.

5.5.14 Ganges-Brahmaputra basin

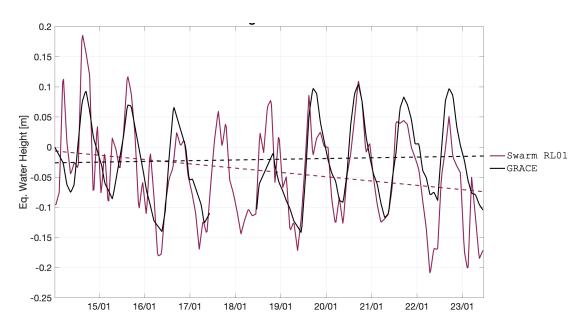


Figure 26 – Time series of EWH for the Ganges-Brahmaputra basin (latitude 15 to 30 degrees, longitude 72 to 89 degrees).

solution	constant	constant	linear term	linear term	corr. coeff.
Solution	term [cm]	term Δ [cm]	[cm/year]	Δ [cm/year]	[]
Swarm RL01	-4.31	-1.39	-0.71	-0.83	0.74
GRACE	-2.92	0.00	0.12	0.00	1.00

Table 16 – Statistics of the agreement between GRACE/GRACE-FO and Swarm time series relative to the GRACE/GRACE-FO climatological model for the Ganges-Brahmaputra basin.

5.5.15 Indochina region

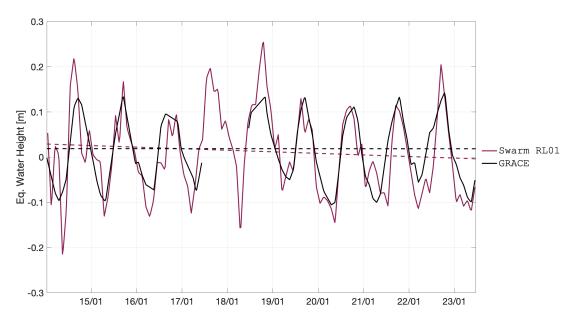


Figure 27 – Time series of EWH for the Indochina region (latitude 12 to 29 degrees, longitude 93 to 105 degrees).

colution	constant	constant	linear term	linear term	corr. coeff.
solution	term [cm]	term Δ [cm]	[cm/year]	Δ [cm/year]	[]
Swarm RL01	0.90	-0.00	-0.35	-0.34	0.80
GRACE	0.90	0.00	-0.00	0.00	1.00

Table 17 – Statistics of the agreement between GRACE/GRACE-FO and Swarm time series relative to the GRACE/GRACE-FO climatological model for the Indochina region.

5.5.16 Northern Australia region

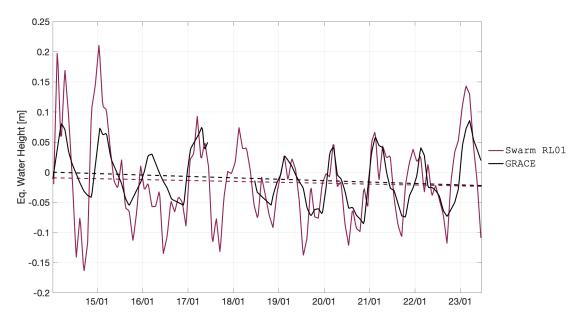


Figure 28 – Time series of EWH for the Northern Australia region (latitude -24 to -10 degrees, longitude 124 to 145 degrees).

solution	constant term [cm]	constant term Δ [cm]	linear term [cm/year]	linear term Δ [cm/year]	corr. coeff.
Swarm RL01	-1.39	-0.69	-0.15	0.09	0.70
GRACE	-0.70	0.00	-0.23	0.00	1.00

Table 18 – Statistics of the agreement between GRACE/GRACE-FO and Swarm time series relative to the GRACE/GRACE-FO climatological model for the Northern Australia region.

5.5.17 Western Antarctica region

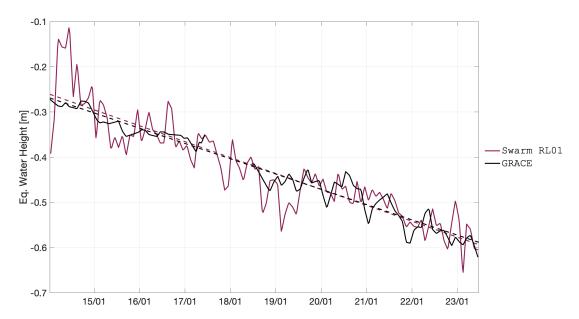


Figure 29 – Time series of EWH for the Western Antarctica region (latitude -80 to -70 degrees, longitude -140 to -85 degrees).

colution	constant	constant	linear term	linear term	corr. coeff.
solution	term [cm]	term Δ [cm]	[cm/year]	Δ [cm/year]	[]
Swarm RL01	-42.71	1.62	-3.52	-0.14	0.91
GRACE	-44.33	0.00	-3.38	0.00	1.00

Table 19 - Statistics of the agreement between GRACE/GRACE-FO and Swarm time series relative to the GRACE/GRACE-FO climatological model for the Western Antarctica region.

5.5.18 Eastern Antarctica region

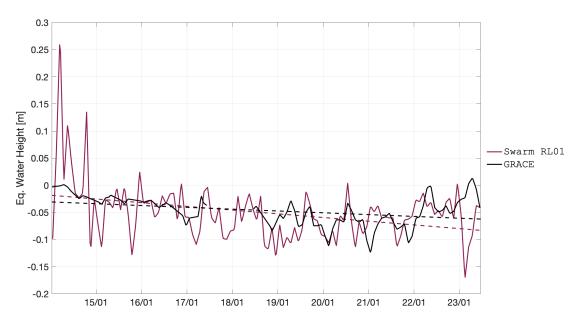


Figure 30 – Time series of EWH for the Eastern Antarctica region (latitude -80 to -68 degrees, longitude 80 to 130 degrees).

solution	constant	constant	linear term	linear term	corr. coeff.
Solution	term [cm]	term Δ [cm]	[cm/year]	Δ [cm/year]	[]
Swarm RL01	-5.05	-0.35	-0.68	-0.35	0.40
GRACE	-4.71	0.00	-0.33	0.00	1.00

Table 20 - Statistics of the agreement between GRACE/GRACE-FO and Swarm time series relative to the GRACE/GRACE-FO climatological model for the Eastern Antarctica region.

2023-10-01 Page 42 of 51

5.5.19 Overview

solution	$\begin{array}{c} \text{constant} \\ \text{term } \Delta \text{ RMS} \\ \text{[cm]} \end{array}$	linear term Δ RMS [cm/year]	corr. coeff. mean []
Swarm RL01	0.90	0.36	0.75
GRACE	0.00	0.00	1.00

Table 21 – Statistics of the agreement between the GRACE and Swarm time series for the regions displayed in Sections 5.5.1 to 5.5.18.

5.6 Temporal variability

temporal STD of Swarm RL01 (2014-01 to 2023-06) 750km Gaussian smoothing

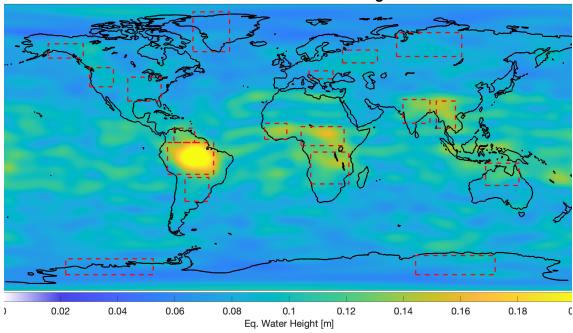


Figure 31 – Temporal variability of the Swarm combined solutions.

A Kinematic Orbits

A.1 Delft University of Technology

Software: GPS High precision Orbit determination Software Tool (GHOST)

(Helleputte, 2004; Wermuth, Montenbruck and Helleputte, 2010)

Preprocessing: None

Differencing Scheme: Undifferenced **Linear combination:** Ionosphere-free

SW_VR_DUT_GS_0018 version 1.0

2023-10-01 Page 43 of 51

GPS observations: Code and carrier phase

Carrier phase ambiguities: Float

Estimator: Bayesian weighted LS

Arc length: 30 hours

Observation weighting: A-priori weights equal to 1m and 1mm for code and phase

observations (resp.)

Data screening: Minimum SNR of 10, minimum of 6 GPS satellites, code and

phase outlier editing threshold of 2 m and 3.5 cm, respectively, 1 meter or larger difference between estimated KO positions

and with Reduced-Dynamic PSO

Transmitter PCV: Official IGS08 ANTEX (Schmid et al., 2007) up to day 17/028, of-

ficial IGS14 ANTEX (Rebischung and Schmid, 2016) afterwards

Receiver PCV: Empirically determined from stacking of reduced-dynamic

POD residuals with 1° binning

GPS orbits and clocks: Final orbits and 5 seconds clocks of CODE (Dach et al., 2017)

Earth precession model: IAU 1976 (Lieske et al., 1977) **Earth nutation model:** IAU 1980 (Seidelmann, 1982)

Earth orientation model: CODE final ERP

A.2 Astronomical Institute of the University of Bern

Software: Bernese v5.3 (Dach et al., 2015)

Preprocessing: Cycle slip detection based on epoch-difference solution

Differencing Scheme: Undifferenced **Linear combination:** Ionosphere-free

Differential code bias: N/A **Ionosphere model:** N/A

GPS observations: Code and carrier phase

Carrier phase ambiguities: Float up to 26 January 2020, ambiguity-fixed afterwards

Estimator: Batch LS
Arc length: 24 hours
Observation weighting: Constant

Data screening: 2 cm/s or larger time-differences of the geometry-free

Transmitter PCV: Official IGS08 ANTEX (Schmid et al., 2007) up to day 17/028, of-

ficial IGS14 ANTEX (Rebischung and Schmid, 2016) afterwards

Receiver PCV: Stacking of carrier phase residuals from reduced-dynamic POD of approx. 120 days, 9 iterations, 1° binning linear combination

of L1B GPS carrier phase observations

GPS orbits and clocks: Final orbits and 5 seconds clocks of CODE (Dach et al., 2017)

Earth precession model: IERS 2010 Conventions (Petit and Luzum, 2010) **Earth nutation model:** IERS 2010 Conventions (Petit and Luzum, 2010)

Earth orientation model: CODE final ERP

A.3 Institute of Geodesy Graz

Software: Gravity Recovery Object Oriented Programming System (GROOPS)

(Mayer-gürr et al., 2020)

SW_VR_DUT_GS_0018 version 1.0

2023-10-01 Page 44 of 51

Preprocessing: Cycle slip detection based on Melbourne-Wuebbena combina-

tion

Differencing Scheme: Raw undifferenced

Linear combination: None (the ionospheric influence is co-estimated)

Differential code bias: Graz University of Technology (TUG) daily estimated absolute

biases

Ionosphere model: Slant total electron content (STEC) 1st, 2nd and 3rd order ef-

fects (Hoque and Jakowski, 2008) estimated in each epoch for

each receiver-transmitter pair

GPS observations: Code and carrier phase

Carrier phase ambiguities: MLAMBDA (Chang, Yang and Zhou, 2005)

Estimator: LS Arc length: 24 hours

Observation weighting: Elevation and azimuth-dependent, ROTI dependent

Data screening: Implicit in VCE

Transmitter PCV: Empirical, estimated from 5.5 years of data, including data

from several LEO missions (GRACE, Jason 2 & 3, MetOp-A & -B, Sentinel 3A, Swarm, TanDEM-X, TerraSAR-X) (Zehentner,

2016)

Receiver PCV: Empirical, spherical harmonics (maximum D/O 100), derived

from 38 months of data

GPS orbits: TUG, estimated using \approx 200 daily IGS stations

GPS clocks: TUG 30 seconds, interpolated using CODE 5 seconds finals to

a sampling of 5 seconds

Earth precession model: IAU 2006/2000A precession-nutation model (Coppola, Seago

and Vallado, 2009)

Earth nutation model: IAU 2006/2000A precession-nutation model (Coppola, Seago

and Vallado, 2009)

Earth orientation model: IERS EOP 08 C04 (Petit and Luzum, 2010)

A.4 Common

Receiver clock corrections: Co-estimated **Phase wind-up:** Correction applied

Sampling rate: 10 seconds up to 15 July 2014, 1 seconds afterwards

Receiver antenna offset: satellite specific values

Elevation cut-off angle: 0°

Swarm attitude: L1B attitude data GPS attitude model: (Kouba, 2009)

B Gravity Field Models

B.1 Astronomical Institute of the University of Bern

Software: Bernese v5.3 (Dach et al., 2015)

Approach: Celestial Mechanics Approach (CMA) (Beutler et al., 2010)

Reference GFM: AIUB-GRACE03S (Jäggi et al., 2011)

Empirical Parameters: Daily and 15 minutes, both piecewise-constant (constrained)

Coord. Axis Combination: TBD

Single Sat. Combination: NEQ, equal weights

SW_VR_DUT_GS_0018 version 1.0

2023-10-01 Page 45 of 51

Temporal correlations: None **Drag Model:** None **EARP and EIRP Models:** None

Non-tidal Model: Unti Nov 2017: AOD1B (Flechtner, Schmidt and Meyer, 2006;

Flechtner, 2007; Flechtner, 2011)

After Nov 2017: AOD1B-RL06 (Dobslaw et al., 2017)

Atmospheric Tidal Model: TBD

Ocean Tidal Model: EOT11a (Savcenko and Bosch, 2012)

Permanent Tide System: tide-free

B.2 Astronomical Institute Ondřejov

Software: (developed in-house)

Approach: Decorrelated Acceleration Approach (DAA) (Bezděk et al., 2014;

Bezděk et al., 2016)

Reference GFM: ITG-GRACE2010s (Mayer-Gürr et al., 2010)

Empirical Parameters: Daily constant-piecewise

Coord. Axis Combination: TBD

Single Sat. Combination: NEQ, equal weights

Temporal correlations: Empirical decorrelation filter

Drag Model: NRLMSISE (Picone et al., 2002)

EARP and EIRP Models: Knocke, Ries and Tapley (1988)

Non-tidal Model: AOD1B-RL06 (Dobslaw et al., 2017)

Atmospheric Tidal Model: Biancale and Bode (2006) **Ocean Tidal Model:** FES2004 (Lyard et al., 2006)

Permanent Tide System: tide-free

B.3 Institute of Geodesy Graz

Software: Gravity Recovery Object Oriented Programming System (GROOPS)

(Mayer-gürr et al., 2020)

Approach: Short-Arcs Approach (SAA) (Mayer-Gürr, 2006)

Reference GFM: GOCO05S (Mayer-Gürr, 2015)

Empirical Parameters: Piecewise linear for each arc (ranging from 15 to 45 minutes)

Coord. Axis Combination: TBD

Single Sat. Combination:

Temporal correlations:

Drag Model:

EARP and EIRP Models:

Non-tidal Model:

NEQ, relative weighting from VCE
Empirical covariance function
JB2008 (Bowman et al., 2008)
Rodriguez-Solano et al. (2012)
AOD1B-RL06 (Dobslaw et al., 2017)

Atmospheric Tidal Model: Biancale and Bode (2006)
Ocean Tidal Model: FES2014 (Carrere et al., 2015)

Permanent Tide System: zero tide

B.4 Ohio State University

Software: (developed in-house)

Approach: Improved Energy Balance Approach (IEBA) (Shang et al., 2015)

Reference GFM: GIF48 (Ries et al., 2011) up to D/O 200

SW_VR_DUT_GS_0018 version 1.0

2023-10-01 Page 46 of 51

Empirical Parameters: 2nd order polynomial every 3 hours, 1-CPR sinusoidal every 24

hours

Coord. Axis Combination: TBD

Single Sat. Combination: NEQ, equal weights

Temporal correlations: None

Drag Model: NRLMSISE (Picone et al., 2002) **EARP and EIRP Models:** Knocke, Ries and Tapley (1988)

Non-tidal Model: AOD1B (Flechtner, Schmidt and Meyer, 2006; Flechtner, 2007;

Flechtner, 2011)

Atmospheric Tidal Model: Biancale and Bode (2006)

Ocean Tidal Model: EOT11a (Savcenko and Bosch, 2012)

Permanent Tide System: tide-free

B.5 Common

Regularization:noneSolid Earth Tidal Model:IERS2010Pole Tidal Model:IERS2010Ocean Pole Tidal Model:IERS2010

Third body perturbations: Sun, Moon, Mercury, Venus, Mars, Jupiter and Saturn, following

the JPL-PLE (Folkner et al., 2014)

 $C_{2,0}$ **coefficient:** estimated alongside other coefficients

Acronyms

AA Acceleration Approach, Rummel (1979)

AIUB Astronomical Institute of the University of Bern, Switzerland,

www.aiub.unibe.ch

AIUB-GRACE03S AIUB GRACE-only static model, version 3, Jäggi et al. (2011)

AOD1B Atmosphere and Ocean De-aliasing Level 1B product, Flechtner, Schmidt and

Meyer (2006), Flechtner (2007) and Flechtner (2011)

AOD1B-RL06 Atmosphere and Ocean De-aliasing Level 1B RL06 product, Dobslaw et al. (2017)

ASU Astronomical Institute (Astronomický ústav), AVCR, Ondřejov,

www.asu.cas.cz/en

AVCR Czech Academy of Sciences (Akademie věd České Republiky), Czech Republic,

www.avcr.cz/en/

CODE Centre for Orbit Determination in Europe, Dach et al. (2017)

CMA Celestial Mechanics Approach, Beutler et al. (2010)

CPR Cycle Per Revolution

CSR Center for Space Research, UT Austin, USA, www.csr.utexas.edu

D/O Degree and Order

DAA Decorrelated Acceleration Approach, Bezděk et al. (2014) and Bezděk et al. (2016)

EARP Earth Albedo Radiation Pressure
EIRP Earth Infrared Radiation Pressure

EBA Energy Balance Approach, O'Keefe (1957) and Jekeli (1999)

EOT Empirical Ocean Tide model

EOT11a 2011 Empirical Ocean Tide model, Savcenko and Bosch (2012)

EWH Equivalent Water Height
EOP Earth Orientation Parameter
ERP Earth Rotation Parameters

SW VR DUT GS 0018 version 1.0

2023-10-01 Page 47 of 51

FES Finite Element Solution global tide model

FES2004 2004 Finite Element Solution global tide model, Lyard et al. (2006) **FES2014** 2014 Finite Element Solution global tide model, Carrere et al. (2015)

Gravity Field Model **GFM**

GIF48 GRACE Intermediate Field 48, Ries et al. (2011)

GNSS Global Navigation Satellite System

Gravity field and steady-state Ocean Circulation Explorer, Balmino et al. (1999) and Floberghagen et al. (2011)**GOCE**

GOCO Gravity Observation COmbination

GOCO05S GOCO release 05 satellite-only gravity field model, Mayer-Gürr (2015)

GPS Global Positioning System

Gravity Recovery And Climate Experiment, Tapley, Reigber and Melbourne (1996) **GRACE**

and Tapley (2004)

GRACE-FO GRACE Follow On, Kornfeld2019 International Astronomical Union IAU

IEBA Improved Energy Balance Approach, Shang et al. (2015)

IERS International Earth Rotation Service

IERS2010 IERS Conventions 2010, Petit and Luzum (2010)

IfG Institute of Geodesy, TUG, Graz, www.ifg.tugraz.at International GNSS Service, Dow, Neilan and Gendt (2005) **IGS** ITG Institut für Geodäsie und Geoinformation, Germany

ITG-GRACE2010s ITG GRACE-only static model, 2010, Mayer-Gürr et al. (2010)

JB2008 Jacchia-Bowman 2008, Bowman et al. (2008)

JPL Jet Propulsion Laboratory, USA, www.jpl.nasa.gov JPL-PLE JPL Planetary and Lunar Ephemerides, Folkner et al. (2014)

KO Kinematic Orbit L₁B Level 1B data

LAMBDA Least-squares Ambiguity De-correlation Adjustment, Teunissen (1995)

LEO Low-Earth Orbit LS least-squares

MLAMBDA Modified LAMBDA method, Chang, Yang and Zhou (2005)

N/A Not Applicable **NEO** Normal Equation

US Naval Research Laboratory Mass Spectrometer and Incoherent Scatter radar atmospheric model, Picone et al. (2002) **NRLMSISE**

OSU Ohio State University, www.osu.edu

PCV Phase Center Variation **POD Precise Orbit Determination**

Precise or Post-processed Science Orbit **PSO**

RL06 Release 6

ROTI Rate of TEC Index **RMS** Root Mean Squared

Short-Arcs Approach, Mayer-Gürr (2006) SAA

SH Spherical Harmonic

Satellite Laser Ranging, Smith and Turcotte (1993) and Combrinck (2010) SLR

SNR Signal-to-Noise Ratio TEC **Total Electron Content**

TU Delft Delft University of Technology, Netherlands, www.tudelft.nl

TUG Graz University of Technology, Austria, www.tugraz.at

SW_VR_DUT_GS_0018 version 1.0

2023-10-01 Page 48 of 51

UT Austin University of Texas at Austin, www.utexas.edu

USA United States of America

VCE Variance Component Estimation

WP Work Package

Symbols

C Stokes coefficient.

References

- Balmino, G. et al. (1999). **The Four Candidate Earth Explorer Core Missions Gravity Field and Steady-State Ocean Circulation Mission**. Tech. rep. SP- 1233(1). Noordwijk, The Netherlands: European Space Agency (cit. on p. 47).
- Bettadpur, Srinivas et al. (2015). **Evaluation of the GGM05 Mean Earth Gravity Model**. In: *EGU General Assembly Conference Abstracts* 17, p. 4153 (cit. on p. 7).
- Beutler, Gerhard et al. (2010). **The celestial mechanics approach: theoretical foundations**. In: *Journal of Geodesy* 84.10, pp. 605–624. DOI: 10.1007/s00190-010-0401-7 (cit. on pp. 44, 46).
- Bezděk, Aleš et al. (2014). **Gravity field models from kinematic orbits of CHAMP, GRACE and GOCE satellites**. In: *Advances in Space Research* 53.3, pp. 412–429. DOI: 10.1016/j.asr. 2013.11.031 (cit. on pp. 45, 46).
- Bezděk, Aleš et al. (2016). **Time-variable gravity fields derived from GPS tracking of Swarm**. In: *Geophysical Journal International* 205.3, pp. 1665–1669. DOI: 10.1093/gji/ggw094 (cit. on pp. 6, 45, 46).
- Biancale, R. and A. Bode (2006). **Mean annual and seasonal atmospheric tide models based on 3-hourly and 6-hourly ECMWF surface pressure data**. Tech. rep. Potsdam, Germany: Deutsches GeoForschungsZentrum GFZ. DOI: 10.2312/GFZ.b103-06011 (cit. on pp. 45, 46).
- Bowman, Bruce et al. (2008). **A New Empirical Thermospheric Density Model JB2008 Using New Solar and Geomagnetic Indices**. In: *AIAA/AAS Astrodynamics Specialist Conference and Exhibit*. August. Reston, Virigina: American Institute of Aeronautics and Astronautics. DOI: 10.2514/6.2008-6438 (cit. on pp. 45, 47).
- Carrere, L et al. (2015). **FES 2014, a new tidal model on the global ocean with enhanced accuracy in shallow seas and in the Arctic region**. In: *EGU General Assembly*. Vienna, Austria (cit. on pp. 45, 47).
- Chang, X. W., Xiaohua Yang and Tianyang Zhou (2005). **MLAMBDA: a modified LAMBDA method for integer least-squares estimation**. In: *Journal of Geodesy* 79.9, pp. 552–565. DOI: 10.1007/s00190-005-0004-x (cit. on pp. 44, 47).
- Cheng, M. and John Ries (2018). **GRACE Technical Note 11: Monthly estimates of C20 from 5 satellites based on GRACE RL06 models**. Austin, USA. URL: https://podaac-tools.jpl.nasa.gov/drive/files/allData/grace/docs/TN-11_C20_SLR.txt (cit. on pp. 2, 7).
- Cheng, Minkang and John Ries (2019). **UT/CSR monthly C20 RL-06 time series from SLR**. Austin, USA. URL: https://filedrop.csr.utexas.edu/pub/slr/degree_2/C20_RL06.txt (cit. on pp. 2, 7).

- Combrinck, Ludwig (2010). **Satellite Laser Ranging**. In: *Sciences of Geodesy I*. Ed. by Guochang Xu. Berlin, Heidelberg: Springer Berlin Heidelberg. Chap. 9, pp. 301–338. DOI: 10.1007/978-3-642-11741-1_9 (cit. on p. 47).
- Coppola, Vincent, John H Seago and David A Vallado (2009). The IAU 2000a and IAU 2006 Precession-Nutation theories and their implementations. In: Advances in the Astronautical Sciences 09.159, pp. 1–20. URL: https://www.agi.com/getmedia/c85a440a-cf71-4e08-ad78-fa73736cee6c/Precession-nutation-Theories-and-their-Implementation.pdf (cit. on p. 44).
- Dach, Rolf et al. (2015). **Bernese GNSS Software Version 5.2**. Bern: Bern Open Publishing. DOI: 10.7892/boris.72297 (cit. on pp. 43, 44).
- Dach, Rolf et al. (2017). **CODE final product series for the IGS**. Bern, Switzerland. DOI: 10. 7892/boris.75876.2 (cit. on pp. 43, 46).
- Dobslaw, H. et al. (2017). A new high-resolution model of non-tidal atmosphere and ocean mass variability for de-aliasing of satellite gravity observations: AOD1B RL06. In: *Geophysical Journal International* 211.1, pp. 263–269. DOI: 10.1093/gji/ggx302 (cit. on pp. 45, 46).
- Dow, J.M. M., R.E. E. Neilan and G. Gendt (2005). **The International GPS Service: Celebrating the 10th anniversary and looking to the next decade**. In: *Advances in Space Research* 36.3, pp. 320–326. DOI: 10.1016/j.asr.2005.05.125 (cit. on p. 47).
- Flechtner, Frank (2007). **Gravity Recovery and Climate Experiment AOD1B Product Description Document for product releases 01 to 04**. Technical report GR-GFZ-AOD-0001. Potsdam: GeoForschungszentrum. URL: ftp://podaac.jpl.nasa.gov/pub/grace/doc/AOD1B_20070413.pdfhttps://www.gfz-potsdam.de/fileadmin/gfz/sec12/pdf/GRACE/AOD1B/AOD1B_20070413.pdf (cit. on pp. 45, 46).
- (2011). GRACE AOD1B RL04 Quality Assurance. Miscellaneous. URL: http://op.gfz-potsdam.de/grace/results/grav/g007_aod1b_rl04.html (visited on 23/07/2015) (cit. on pp. 45, 46).
- Flechtner, Frank, Roland Schmidt and Ulrich Meyer (2006). **De-aliasing of Short-term Atmospheric and Oceanic Mass Variations for GRACE**. In: *Observation of the Earth System from Space*. Ed. by J. Flury et al. Springer Berlin Heidelberg, pp. 83–97. DOI: 10.1007/3-540-29522-4_7 (cit. on pp. 45, 46).
- Floberghagen, Rune et al. (2011). **Mission design, operation and exploitation of the gravity field and steady-state ocean circulation explorer mission**. In: *Journal of Geodesy* 85.11, pp. 749–758. DOI: 10.1007/s00190-011-0498-3 (cit. on p. 47).
- Folkner, William M et al. (2014). **The Planetary and Lunar Ephemerides DE430 and DE431**. In: *Interplanet. Netw. Prog. Rep* 42.196. URL: https://ipnpr.jpl.nasa.gov/progress_report/42-196/196C.pdf (cit. on pp. 46, 47).
- Guo, J. Y., X. J. Duan and C. K. Shum (2010). **Non-isotropic Gaussian smoothing and leakage reduction for determining mass changes over land and ocean using GRACE data**. In: *Geophysical Journal International* 181.1, pp. 290–302. DOI: 10.1111/j.1365-246X.2010.04534.x (cit. on p. 7).
- Guo, J. Y. et al. (2015). **On the energy integral formulation of gravitational potential differences from satellite-to-satellite tracking**. In: *Celestial Mechanics and Dynamical Astronomy* 121.4, pp. 415–429. DOI: 10.1007/s10569-015-9610-y (cit. on p. 6).
- Helleputte, T. van (2004). **GPS High Precision Orbit Determination Software Tools: User Manual**. Oberpfaffenhofen (cit. on p. 42).
- Hoque, M. Mainul and N. Jakowski (2008). **Estimate of higher order ionospheric errors in GNSS positioning**. In: *Radio Science* 43.5, n/a–n/a. DOI: 10.1029/2007RS003817 (cit. on p. 44).

- Jäggi, A. et al. (2011). **AIUB-GRACE03S**. Bern, Switzerland. URL: http://icgem.gfz-potsdam.de/(cit. on pp. 44, 46).
- Jäggi, A. et al. (2016). **Swarm kinematic orbits and gravity fields from 18 months of GPS data**. In: *Advances in Space Research* 57.1, pp. 218–233. DOI: 10.1016/j.asr.2015.10.035 (cit. on p. 6).
- Jekeli, Christopher (1999). **The determination of gravitational potential differences from satellite-to-satellite tracking**. In: *Celestial Mechanics and Dynamical Astronomy* 75.2, pp. 85–101. DOI: 10.1023/A:1008313405488 (cit. on p. 46).
- Knocke, P., J. Ries and B. Tapley (1988). **Earth radiation pressure effects on satellites**. In: *Astrodynamics Conference*. Reston, Virigina: American Institute of Aeronautics and Astronautics. DOI: 10.2514/6.1988-4292 (cit. on pp. 45, 46).
- Kouba, J. (2009). **A simplified yaw-attitude model for eclipsing GPS satellites**. In: *GPS Solutions* 13.1, pp. 1–12. DOI: 10.1007/s10291-008-0092-1 (cit. on p. 44).
- Lieske, J H et al. (1977). Expression for the precession quantities based upon the IAU (1976) system of astronomical constants. In: *Astronomy and Astrophysics* 58, pp. 1–16 (cit. on p. 43).
- Loomis, B.D. and K.E. Rachlin (2020). **GRACE Technical Note 14: NASA GSFC SLR C20 and C30 solutions**. Greenbelt, USA. URL: https://podaac-tools.jpl.nasa.gov/drive/files/allData/gracefo/docs/TN-14_C30_C20_GSFC_SLR.txt (cit. on pp. 2, 7).
- Lyard, Florent et al. (2006). **Modelling the global ocean tides: modern insights from FES2004**. In: *Ocean Dynamics* 56.5-6, pp. 394–415. DOI: 10.1007/s10236-006-0086-x (cit. on pp. 45, 47).
- Mayer-Gürr, Torsten (2006). **Gravitationsfeldbestimmung aus der Analyse kurzer Bahnbögen am Beispiel der Satellitenmissionen CHAMP und GRACE**. Doctoral dissertation. Rheinischen Friedrich-Wilhelms Universität Bonn. URL: http://hss.ulb.uni-bonn.de/2006/0904/0904.pdf (cit. on pp. 45, 47).
- (2015). **The Combined Satellite Gravity Field Model GOCO05s**. In: *EGU General Assembly*. EGU2015-12364. Vienna, Austria (cit. on pp. 45, 47).
- Mayer-Gürr, Torsten et al. (2010). **ITG-Grace2010: the new GRACE gravity field release computed in Bonn**. In: *EGU General Assembly*. EGU2010-2446. Vienna, Austria. URL: http://www.igg.uni-bonn.de/apmg/index.php?id=itg-grace2010 (cit. on pp. 45, 47).
- Mayer-gürr, Torsten et al. (2020). **GROOPS: A software toolkit for gravity field recovery and GNSS processing**. In: *Earth and Space Science Open Archive*. DOI: 10.1002/essoar. 10505041.1 (cit. on pp. 43, 45).
- O'Keefe, John A. (1957). **An application of Jacobi's integral to the motion of an earth satellite**. In: *The Astronomical Journal* 62, p. 265. DOI: 10.1086/107530 (cit. on p. 46).
- Petit, Gérard Gerard and Brian Luzum (2010). **IERS Conventions (2010)**. Frankfurt am Main. URL: http://www.iers.org/TN36/ (cit. on pp. 43, 44, 47).
- Picone, J. M. et al. (2002). **NRLMSISE-00 empirical model of the atmosphere: Statistical comparisons and scientific issues**. In: *Journal of Geophysical Research: Space Physics* 107.A12, SIA 15–1–SIA 15–16. DOI: 10.1029/2002JA009430 (cit. on pp. 45–47).
- Rebischung, P and R Schmid (2016). **IGS14/igs14.atx: a new Framework for the IGS Products**. In: *AGU Fall Meeting Abstracts*. Vol. 2016, G41A–0998 (cit. on p. 43).
- Ries, John C. et al. (2011). **Mean Background Gravity Fields for GRACE processing**. In: *GRACE Science Team Meeting*. Austin, USA. URL: http://download.csr.utexas.edu/pub/grace/Proceedings/Presentations_GSTM2011.pdf (cit. on pp. 45, 47).
- Rodriguez-Solano, C. J. et al. (2012). **Impact of Earth radiation pressure on GPS position estimates**. In: *Journal of Geodesy* 86.5, pp. 309–317. DOI: 10.1007/s00190-011-0517-4 (cit. on p. 45).

- Rummel, R. (1979). **Determination of short-wavelength components of the gravity field from satellite-to-satellite tracking or satellite gradiometry**. In: *Manuscripta Geodaetica* 4.2, pp. 107–148 (cit. on p. 46).
- Savcenko, R and W Bosch (2012). **EOT11a Empirical ocean tide model from multi-mission satellite altimetry**. Tech. rep. München, Germany: Deutsches Geodätisches Forschungsinstitut. URL: https://epic.awi.de/36001/1/DGFI_Report_89.pdf (cit. on pp. 45, 46).
- Schmid, Ralf et al. (2007). **Generation of a consistent absolute phase-center correction model for GPS receiver and satellite antennas**. In: *Journal of Geodesy* 81.12, pp. 781–798. DOI: 10.1007/s00190-007-0148-y (cit. on p. 43).
- Seidelmann, P. K. (1982). **1980 IAU Theory of Nutation: The final report of the IAU Working Group on Nutation**. In: *Celestial Mechanics* 27.1, pp. 79–106. DOI: 10.1007/BF01228952 (cit. on p. 43).
- Shang, Kun et al. (2015). **GRACE time-variable gravity field recovery using an improved energy balance approach**. In: *Geophysical Journal International* 203.3, pp. 1773–1786. DOI: 10.1093/gji/ggv392 (cit. on pp. 45, 47).
- Smith, David E. and Donald L. Turcotte (1993). **Millimeter Accuracy Satellite Laser Ranging: a Review**. In: *Contributions of Space Geodesy to Geodynamics: Technology*. Ed. by John J. Degnan. Vol. 25. Geodynamics Series. Washington, D. C.: American Geophysical Union. DOI: 10.1029/GD025p0133 (cit. on p. 47).
- Suesser-Rechberger, Barbara et al. (2022). **Improved precise kinematic LEO orbits based on the raw observation approach**. In: *Advances in Space Research*. DOI: 10.1016/j.asr. 2022.03.014 (cit. on p. 6).
- Tapley, B., C. Reigber and W Melbourne (1996). **Gravity Recovery And Climate Experiment** (**GRACE**) mission. Baltimore, USA (cit. on p. 47).
- Tapley, Byron D. (2004). **GRACE Measurements of Mass Variability in the Earth System**. In: *Science* 305.5683, pp. 503–505. DOI: 10.1126/science.1099192 (cit. on p. 47).
- Teixeira da Encarnação, João and Pieter Visser (2017). **TN-01 : Standards and Background Models**. Tech. rep. Delft, the Netherlands: Delft University of Technology. DOI: 10.13140/RG.2.2.12840.32006/1 (cit. on p. 6).
- (2019). TN-03: Swarm models validation. Tech. rep. TU Delft. DOI: 10.13140/RG.2.2.
 33313.76640 (cit. on p. 6).
- Teunissen, P. J. G. (1995). **The least-squares ambiguity decorrelation adjustment: a method for fast GPS integer ambiguity estimation**. In: *Journal of Geodesy* 70.1-2, pp. 65–82. DOI: 10.1007/BF00863419 (cit. on p. 47).
- Wermuth, Martin, Oliver Montenbruck and Tom Van Helleputte (2010). **GPS high precision orbit determination software tools (GHOST)**. In: *4th International Conference on Astrodynamics Tools and Techniques*. Madrid: ESA WPP-308 (cit. on p. 42).
- Zehentner, Norbert (2016). **Kinematic orbit positioning applying the raw observation approach to observe time variable gravity**. Doctoral dissertation. Graz University of Technology, p. 175. DOI: 10.13140/RG.2.2.33916.33927 (cit. on p. 44).